NOAFAULTS: A DIGITAL DATABASE FOR ACTIVE FAULTS IN GREECE

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Abstract

This paper documents the approach to compiling a digital database of fault geometry and additional attributes primarily to support seismicity monitoring at the National Observatory of Athens (NOA). A database for Greek active faults has been constructed from published fault maps in peer-reviewed journals since 1972. The standard commercial software ARC GIS has been used to design and populate the database. The fault layer was produced by on-screen digitization and is available to the scientific community in ESRI shapefile (SHP) and TXT formats in WGS84 projection. A KML file is also available to display the fault data in an Earth browser such as Google Earth. In this version of the database, we focus our attention to the active faults of the upper (Aegean + Eurasian) plate and the back-arc region of the Hellenic Arc, in general. 963 faults are included. The database is freely accessible from the Internet.

Key words: Active faults, GIS, Hellenic Arc.

1. Introduction

Greece is well known in the geosciences for its active tectonics and accompanying seismicity (e.g. Skourphios, 1894; Maravelakis, 1943; Papazachos and Delibasis, 1969; McKenzie, 1978; Delibasis et al., 1981; Mariolakos et al., 1981; Mercier et al., 1983; Pavlides and Mountrakis, 1987; Mariolakos et al., 1989; Ambraseys and Jackson, 1990; Roberts and Jackson, 1991; Taymaz et al., 1991; Doutsos and Poulimenos, 1992; Armijo et al., 1992; Caputo and Pavlides, 1993; Jackson, 1994; Roberts and Ganas, 2000; Lekkas, 2001; Goldsworthy et al., 2002; Caputo et al., 2010; Figure 1). A search on the ISI Web of knowledge database (during March 2013) for Greece AND active faults produced 458 papers for the period 1970-present. Almost every single paper refers to
an active fault, providing some information on its geometry, kinematics and/or earthquake potential. Yet, a complete active fault database in digital form is missing for this country.

Figure 1 - Relief map of the Hellenic Arc and the Aegean Sea. The surface trace of the Hellenic Arc is shown as adapted after Kreemer and Chamot-Rooke (2004). Also shown are large faults – plate boundaries (North Anatolian Fault - CTF stands for Cephalonia Transform Fault), GPS vectors and selected earthquake focal mechanisms from the Global CMT database http://www.globalcmt.org/ (compressional quadrants are shown in green). Image after Ganas and Parsons (2009).

We present first results of a project which stems from the everlasting need to construct seismicity maps containing accurate traces of faults (neotectonic and/or active) for routine seismic monitoring at NOA. During co-seismic and early post-seismic phase (24 hrs) there is a growing pressure by scientists, engineers, politicians and the media to associate earthquake epicentres with faults or faultlines. This is because we have identified two main hazards associated with active faults: ground rupture and ground shaking associated with the rupture. While ground rupture affects only the area in the immediate vicinity of the fault, the ground shaking will impact a much larger area. Existing maps in Greece (e.g. Delibasis et al, 1981; IGME, 1989; Papazachos et al., 2001) can be used for synoptic views because of their small scale (1:275000 or smaller). The Earthquake Planning and Protection Organisation of Greece (EPPO; www.oasp.gr) has published a series of 1:100000 neotectonic map sheets (ten sheets as of June 2013), only in hard copy. The
Greek Database of Seismogenic Sources (Gre.Da.S.S.) was presented by Pavlides et al., (2008) and Pavlides et al., (2010), focused on the North Aegean region and by Caputo et al., (2012) and Sboras (2012) on northern Greece.

Recently, large-scale (equal to or greater than 1:50000) maps of active faults are required by the Greek Ministries of Environment and Infrastructure in order for city plan expansion to be approved (e.g. Ganas et al., 2010). Most major cities of Greece like Athens, Thessaloniki, Patras and Heraklion host large ports and airports, refineries, large industrial plants, gas pipelines etc which are infrastructures of high economic importance. According to the current Building Code of Greece (EAK, 2000) and to Eurocode 8 regulations (EC-8, 2002) active faults need to be taken seriously in consideration in city planning and infrastructure projects. For the purposes of this study we consider faults as active if they show geologic evidence for slip during the last 125000-130000 years. As geologic evidence we consider published maps or reports or cross-sections documenting displacement of Holocene – late Pleistocene deposits and/or landforms. A longer time period than the Holocene is more appropriate, especially since most earthquake recurrence intervals along intra-plate faults are thousands to tens of thousands of years (see Crone et al., 1997, and Machette, 2000 for the paleoseismic perspective). Also, this age limit is followed by many countries (USA – WSSPC 1997, AIST Japan, New Zealand etc; see reference section for government sources). We also include offshore faults with seismological evidence for their activity during instrumental times (e.g. the 2001 Skyros earthquake fault, Benetatos et al., 2002).

In terms of active tectonics in Greek territory and its boundaries we may distinguish two major settings: the Hellenic Arc and the Greek mainland including the Aegean Sea and parts of western Anatolia (Figure 1). The Hellenic Arc defines a plate boundary between Eurasia and Africa lithospheric plates (Nubia; Ganas and Parsons, 2009 and references there in). We will not consider faults or fault segments along the plate interface as no consensus neither exist on the extent of past ruptures or on the degree of seismic coupling (see papers by Laigle et al., 2002, 2004; Ganas et al., 2009; Shaw and Jackson, 2010) along the arc. Therefore, we focus this first version of the database to the upper plate. The upper plate comprises the provinces of Central Greece, Macedonia, Thrace, Epirus, Peloponnese, Cyclades, Crete and Dodecanese. The latter two lie in a prominent position in the fore-arc of the Hellenic Subduction Zone where large, shallow earthquakes are most likely to occur by normal and strike-slip faulting (e.g. Delibasis et al., 1981; Armijo et al., 1992; Ganas and Parsons, 2009; Caputo et al., 2010). The upper crust of the fore-arc region displays numerous evidence of extensional deformation along both arc-parallel, high-angle, E-W striking faults and along arc-normal, high-angle pure-normal and oblique-normal faults that strike on average N-S. Other sets of normal faults striking NW-SE and NE-SW occur as well. The Greek mainland also deforms by normal faulting (e.g. Ambraseys and Jackson, 1990; Roberts and Jackson, 1991; Doutsos and Poulimenos, 1992; Piccardi, 2000; Kokkalas et al., 2007) while strike-slip tectonics prevail in the central and north Aegean Sea (Taymaz et al., 1991; Koukouvelas and Aydin, 2002).

2. Material and Methods

2.1. Identification of Faults

The digitized faults were identified and selected from literature sources, mainly published papers in peer-review journals. These papers were published during the period 1970-2010 and contain fault maps of active faults at a variety of scales. The majority of faults show normal-slip kinematics. The parameterization of the faults includes typical elements such as geometry, kinematics, evidence for activity etc (Table 1). The geological setting of the fault was determined according to Papanikolaou (1989) in most cases. The fault attributes were inserted into an Excel spreadsheet containing the main characteristics of each structure (geometry, kinematics) and were imported to the ARCGIS v 9.3. Number of active faults included in the database: 963.

To design and parameterize the fault database we followed the template produced under the EU project COST 625 (2000-2006; http://fir.seismology.hu/cost625/index.html )
2.2. Method of Work

The working method comprises the following stages: a) scanning of the fault map included in the selected publication in TIFF format (e.g. Figure 2 showing fault traces in north Parnitha region, central Greece; Figure 3 shows a field photograph of the fault surface) b) georeferencing of the scanned raster (TIFF) image using ArcGIS v9.3 software c) digitization of the fault traces as ESRI shapefiles (SHP) d) editing of the attribute table using published information such as fault length, dip direction, dip angle, slip rate etc (Figure 4).

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Type</th>
<th>Description</th>
</tr>
</thead>
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<tr>
<td>Fault code</td>
<td>Number</td>
<td>Fault number</td>
</tr>
<tr>
<td>Fault name</td>
<td>Text</td>
<td>Name of structure</td>
</tr>
<tr>
<td>Kinematics</td>
<td>Text</td>
<td>Normal/Reverse/Strike slip sinistral /Strike slip dextral</td>
</tr>
<tr>
<td>Geological Setting</td>
<td>Text</td>
<td>Isopic zone of the Hellenides</td>
</tr>
<tr>
<td>Location Reliability</td>
<td>Text</td>
<td>At mapping scale: (1:50000 and above) is precise. (1:50000&lt;scale&lt;1:250000) is approximate. (1:250000 and below) or map scale taken from seismological data is inferred.</td>
</tr>
<tr>
<td>Monitoring</td>
<td>Text</td>
<td>Description of possible monitoring networks.</td>
</tr>
<tr>
<td>Paleoseismology</td>
<td>Text</td>
<td>Description of possible paleoseismological studies.</td>
</tr>
<tr>
<td>Creep</td>
<td>Text</td>
<td>Yes/No</td>
</tr>
<tr>
<td>Co-seismic slip (m)</td>
<td>Text</td>
<td>Yes/No</td>
</tr>
<tr>
<td>Notes</td>
<td>Text</td>
<td>If necessary</td>
</tr>
<tr>
<td>Activity reliability</td>
<td>Text</td>
<td>Proven (from trench or seismic rupture) / Very Likely (from geomorphological criteria, or microseismic activity) / Possible (from orientation of structure with respect to the current stress field).</td>
</tr>
<tr>
<td>Slip rate (mm/yr)</td>
<td>Number</td>
<td>If known</td>
</tr>
<tr>
<td>Instrumental Seismicity</td>
<td>Text</td>
<td>Yes/No</td>
</tr>
<tr>
<td>Historical Seismicity</td>
<td>Number</td>
<td>Yes/No</td>
</tr>
<tr>
<td>Seismic Events</td>
<td>Text</td>
<td>List of earthquakes attributed to this fault</td>
</tr>
<tr>
<td>Max Magnitude</td>
<td>Number</td>
<td>Maximum recorded magnitude (Ms or Mw)</td>
</tr>
<tr>
<td>Risk level</td>
<td>Text</td>
<td>Crossing or in the vicinity (5 km) of towns - hazardous facilities (refineries etc). Low if fault length is less than 1 km. Medium if 1km &lt;fault length &lt;5km. High if fault length is more than 5 km.</td>
</tr>
<tr>
<td>Study quality</td>
<td>Text</td>
<td>Detailed/To be improved</td>
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<tr>
<td>Notes</td>
<td>Text</td>
<td>If necessary</td>
</tr>
<tr>
<td>Figure</td>
<td>Number</td>
<td>Figure with digitized faults (from original paper)</td>
</tr>
<tr>
<td>References</td>
<td>Text</td>
<td>Author(s) and title of paper(s)</td>
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<tr>
<td>Compiler</td>
<td>Text</td>
<td>Name of the compilers</td>
</tr>
</tbody>
</table>

The following control layers were used i) a vector (SHP format; standard industry format) file of the coastline of Greece at the scale of 1:250000 and ii) the longitude-latitude coordinates provided by the original fault map itself, as published in the literature. The coastline layer had the Greek national projection (Photogrammetric Engineering and Remote Sensing, 2002). All maps based on the coastline control layer were re-projected to geographic coordinates in WGS84. In the second case (based on the original map), we assumed that their geographic coordinates represent tick points under the WGS 1984 ellipsoid. The error of the georeferencing is a function of 1) the polygon used in the rectification of each map 2) the original map scale and 3) the quality of the con-
trol layer. Depending on the scale of the original map we calculated registration errors at the range 12-450 m (the latter error refers to offshore faults).

The digitization of faults took place on-screen using appropriate zooming, paying attention to all fault geometry details provided by the authors. Each fault segment is represented by a vector feature of the database. Then, characteristics of the fault segment such as fault strike, dip, rake, length, slip rate, co-seismic displacement etc were added to the appropriate fields in the database. Individual SHP files were converted to TXT (ASCII) files and then were grouped in 1 set to facilitate further processing.

Figure 2 - Geological map showing the Maliza active fault in central Greece (after Ganas et al., 2004). Stereonets show results from slip data collection localities.

Figure 3 - Field photograph of the Maliza active fault (locality N5 on the map in Fig. 2). View to the south.
The database contains no information on the 3-D shape of faults. All faults are shown in 2-D as lines (Figure 5). We assume that the finite shape of normal and reverse faults is elliptical and that of strike-slip faults is rectangular, however, the user of the database should not make any assumption on the fault shape of the feature in study. Details on fault geometry and kinematics should be looked at original papers.

The SHP files were also converted to TXT format so that GMT software users (Wessel and Smith, 1995) could access the dataset and plot it on their maps. In this case, each fault is represented by one TXT file containing the geographic coordinates of start-end and intermediate vertices. No fault attributes are included. A KML file is also available to display the fault data in an Earth browser such as Google Earth http://earth.google.com / (Figure 6). All three file formats are available from www.gein.noa.gr

3. Discussion

3.1. Ambiguity of Fault Location on Fault Maps

In Greece, topographic maps for general use are available at 1:50000 and 1:100000 scales from the Hellenic Military Geographical Service (HMGS). They use a national projection system called EGS/A’87 (Photogrammetric Engineering and Remote Sensing, 2002), which is a Transverse Mercator Projection, mapping Greece in one zone. A few areas are also available at 1:25000 scale. Most geologists use such maps as cartographic background to draw their faults. Following this reality, we introduced the following criteria to differentiate among fault location reliability.

Precise fault trace: the map from which the fault was digitized is 1:50000 scale or larger (e.g. Papanikolau et al., 1998; Morewood and Roberts, 1999, 2001, 2002; Ganas et al., 2004). For Greece, the 1:50000 maps carry an accuracy of 15 m in the horizontal and 5 m on the vertical.

Approximate fault trace: the map from which the fault was digitized is 1:50000 scale or smaller (up to 1:250000; e.g. Jenkins, 1972; Mettos et al., 1991; King et al., 1993; Galanakis et al., 1998). Also, in case of no surface expression the location can be approximated by good-quality seismological data where epicentre planimetric accuracy is better than 2 km (e.g. seismic fault of the 2001 earthquake, offshore Skyros, Benetatos et al., 2002, Figure 4; seismic fault of the 8 June 2008 earthquake in NW Peloponnese; Ganas et al., 2009).

Inferred fault trace: the location of the fault is based on published seismological data (epicentres of aftershocks) of regional networks or seen in offshore seismic profiles (e.g. Laigle et al., 2000; Stefatos et al., 2002; Ganas et al., 2013).
3.2. Ambiguity on Fault Activity Reliability

We introduced the following criteria to differentiate among fault activity reliability:

Proven activity as demonstrated by reported co-seismic displacement along the fault trace during large earthquakes (e.g. Pavlides and Tranos, 1991; Pavlides et al., 1995; Ambraseys and Jackson, 1998), including the age of last motion measured in paleoseismological trenches (e.g. Heliki fault in Peloponnese, Koukouvelas et al., 2001; Atalanti fault in central Greece; Pantosti et al., 2004; Kaparelli and Mygdonia faults in Chatzipetros et al., 2005). Other evidence of present activity includes displacement of Holocene-Late Pleistocene sedimentary deposits and/or Middle-Late
Pleistocene sedimentary deposits (e.g. Caputo, 1996; Ganas, 1997; Maroukian et al., 2000; Goldsworthy and Jackson, 2001).

Activity Very likely as demonstrated by a) microseismic activity aligned for several kilometres for half- to two fault lengths as seen on seismicity maps (e.g. Louvari et al., 1999; Ganas et al., 2009), b) geological indicators (bean-shaped basins filled with Neogene syn-rift etc; e.g. Roberts and Ganas, 2000) and c) geomorphologic indicators (systematically displaced streams and erosion surfaces, existence of triangular facets along mountain fronts, etc; Papanikolaou et al., 1988; Ganas, 1997; Goldsworthy and Jackson, 2000).

Activity Possible as demonstrated by the orientation (strike) of fault with respect to the current stress field (e.g. offshore faults along the Hellenic Arc; Masele et al., 1982; Hatzfeld et al., 1990), and/or by faults located near sharp discontinuities on GPS profiles (e.g. faults reported by Underhill, 1988, onshore Zakynthos where a GPS profile discontinuity exists in Hollenstein et al., 2008).

3.3. Overlapping Sources of Information

Many active faults are included in overlapping maps of the same area by different authors. For example the ISI Web of Knowledge database contains 149 papers on active faults AND Gulf of Corinth (June 2013 search). In such cases, we proceeded with a careful selection of those papers that contain the original fault maps, both onshore and offshore. It is possible that we missed a few
papers in this research as we could not trace back in literature all published papers. We ask for the understanding of those authors that see their work not mentioned in this database. As we present this product in its preliminary version, we aim to integrate this knowledge in future editions of this database. If possible, we also ask for their inputs by sending original material to the first author in Athens.

3.4. Lack of Consensus

There were cases where the fault trace is debatable among various authors, so as a decision had to be made by us. For example, old maps of the Atalanti fault in central Greece (IGME, 1989) show this active fault to reach the town of Agios Konstantinos (about 15 km to the northwest of Atalanti) while newer maps (Ganas, 1997; Ganas et al., 1998; Cundy et al., 2000) show it to terminate near the town of Atalanti. It is unfortunate that this version of the database cannot present all published versions of Greek active faults. Therefore, a disclaimer has been added at the end of this paper cautioning users to probable errors and omissions of this product.

3.5. Future Steps

We plan to update the database at regular intervals. Future add-ons may include details on fault geometry (strike, dip angle and dip direction) and kinematics (rake angle, sense of displacement), field photographs of selected fault planes known for hosting large earthquakes, GPS measurements from permanent stations close to major faults, links to on-line sources (libraries, etc), historical seismicity data etc.

We also anticipate that the database may be used in seismic hazard assessments (assuming the active faults as seismic sources) and as a tool for earthquake risk assessment in various parts of Greece. The database is more complete in the regions of central Greece (including Thessaly and Peloponnese), in west and central Macedonia and along the Hellenic Arc. The Hellenic Arc is prone to tsunamis (e.g. Lorito et al., 2008) so the database may be used in tsunami early warning applications. We could not retrieve lots of literature pieces for eastern Macedonia, Thrace and the Aegean Sea, so these areas will become targets for future updates of the database.

4. Conclusions

This paper documents the approach to compiling a digital database of fault geometry and additional attributes primarily to support seismic operations at the National Observatory of Athens (NOA).

A first version of the digital database of Greek active faults has been constructed. The database contains 963 faults published in peer-reviewed literature. This product may be used by geologists, seismologists and engineers working with cartographic scales 1:50000 and smaller. We exercise extreme caution to users of the database working with larger scales (e.g. 1:10000). The database is a growing database and will be subject to change as new information becomes available and new maps are published on Greek faults.

5. Acknowledgments

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